



**British
Geological Survey**

NATURAL ENVIRONMENT RESEARCH COUNCIL

Applied geoscience for our
changing Earth

celebrating
175
years

Estimating the geomagnetic field as a reference for wellbore surveys: accounting for all sources and uncertainties

Ellen Clarke

The Earth's magnetic field

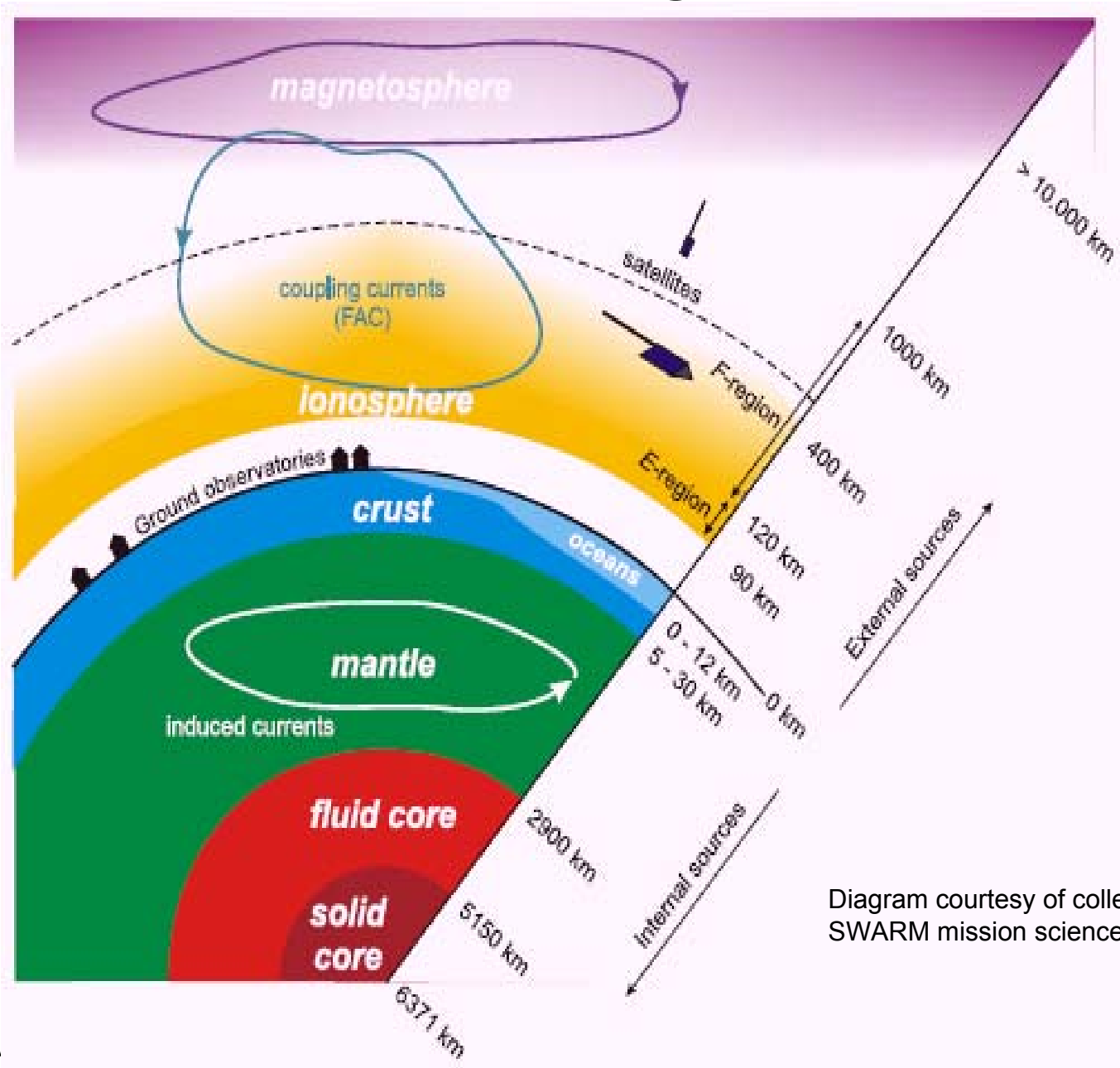


Diagram courtesy of colleagues in the SWARM mission science team



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years

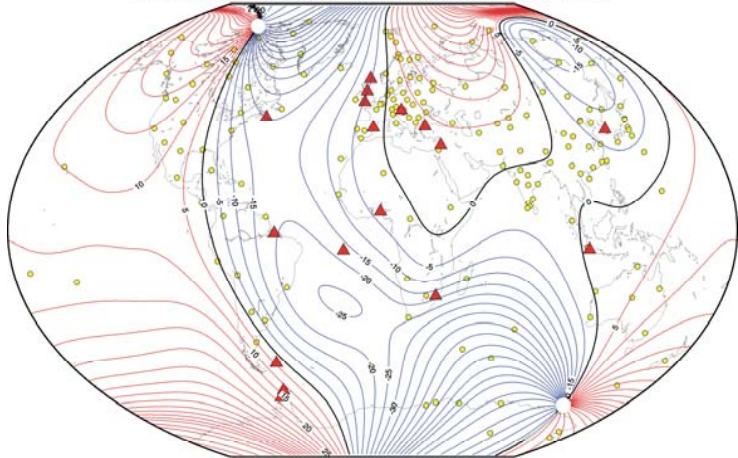
The Earth's magnetic field

- Most of the field is from the **Earth's core**
 - varies slowly with time (**months to years**)
- Local fields from magnetized rocks in **Earth's crust**
 - relatively **stable** with time
- Fields due to currents in the **ionosphere** and the **magnetosphere**
 - variations from **seconds to years**

Geomagnetic field sources ...

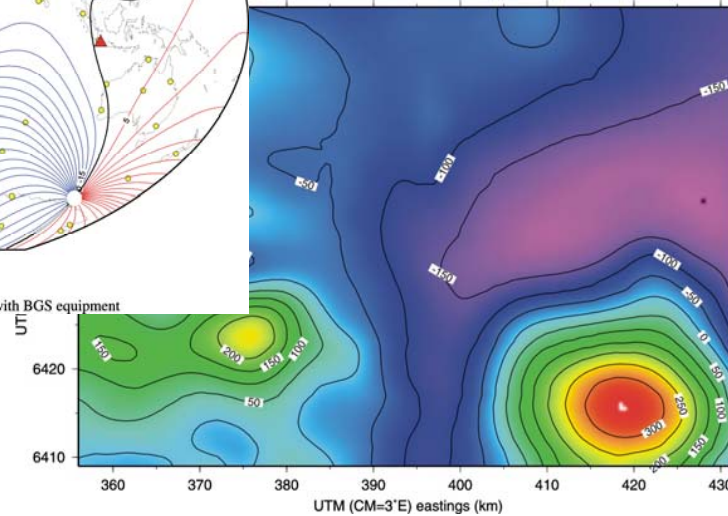
GLOBAL GEOMAGNETIC FIELD MODELLING

Contours are of declination at 2001 (red: degrees east of true north, blue: west)

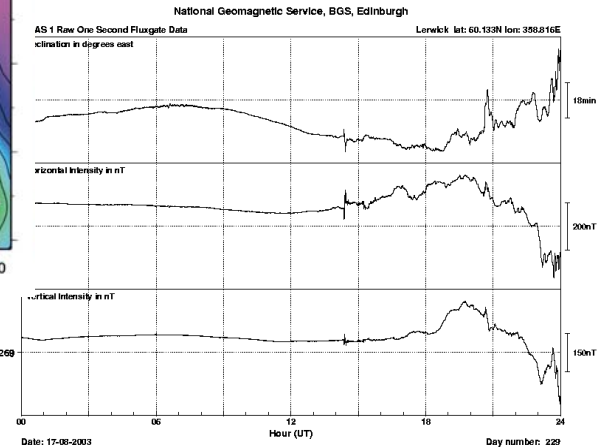


- Geomagnetic observatories
- ▲ Geomagnetic observatories/stations with BGS equipment

Core Field



Crustal Field

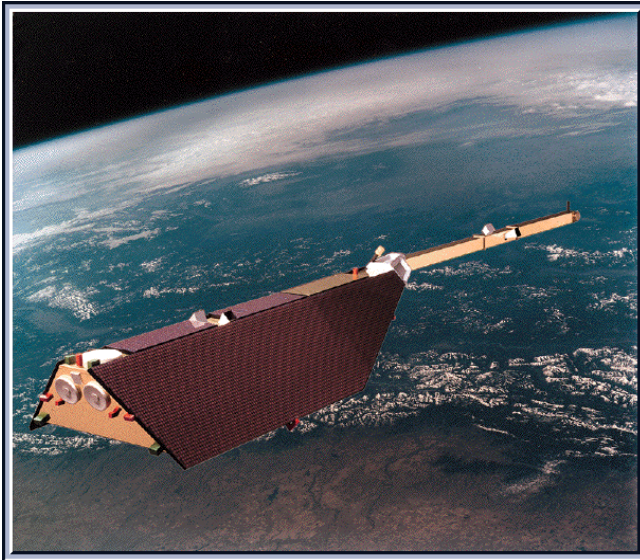


External Fields



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... and how they are measured



Magnetic survey satellite



Aeromagnetic survey



Magnetic Observatory

Directional reference and uncertainty

Reference direction for drilling $\rightarrow D + \varepsilon$

1. Ideally, account for all sources

$$D_1 = D_{core} + D_{crust} + D_{external}$$

$$\varepsilon_1 = \varepsilon_{core} + \varepsilon_{crust} + \varepsilon_{external}$$

2. When external fields are ignored

$$D_2 = D_{core} + D_{crust} + 0$$

$$\varepsilon_2 = \varepsilon_{core} + \varepsilon_{crust} + D_{external}$$

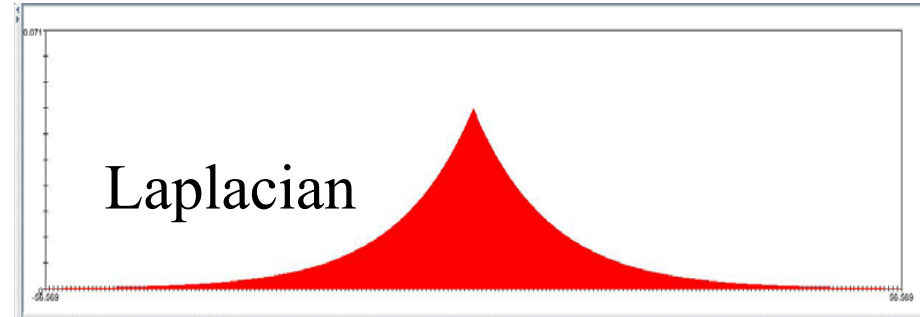
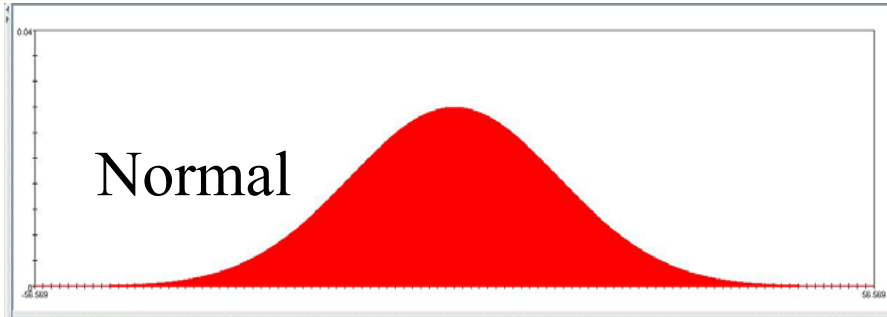
3. When crustal and external fields are ignored

$$D_3 = D_{core} + 0 + 0$$

$$\varepsilon_3 = \varepsilon_{core} + D_{crust} + D_{external}$$

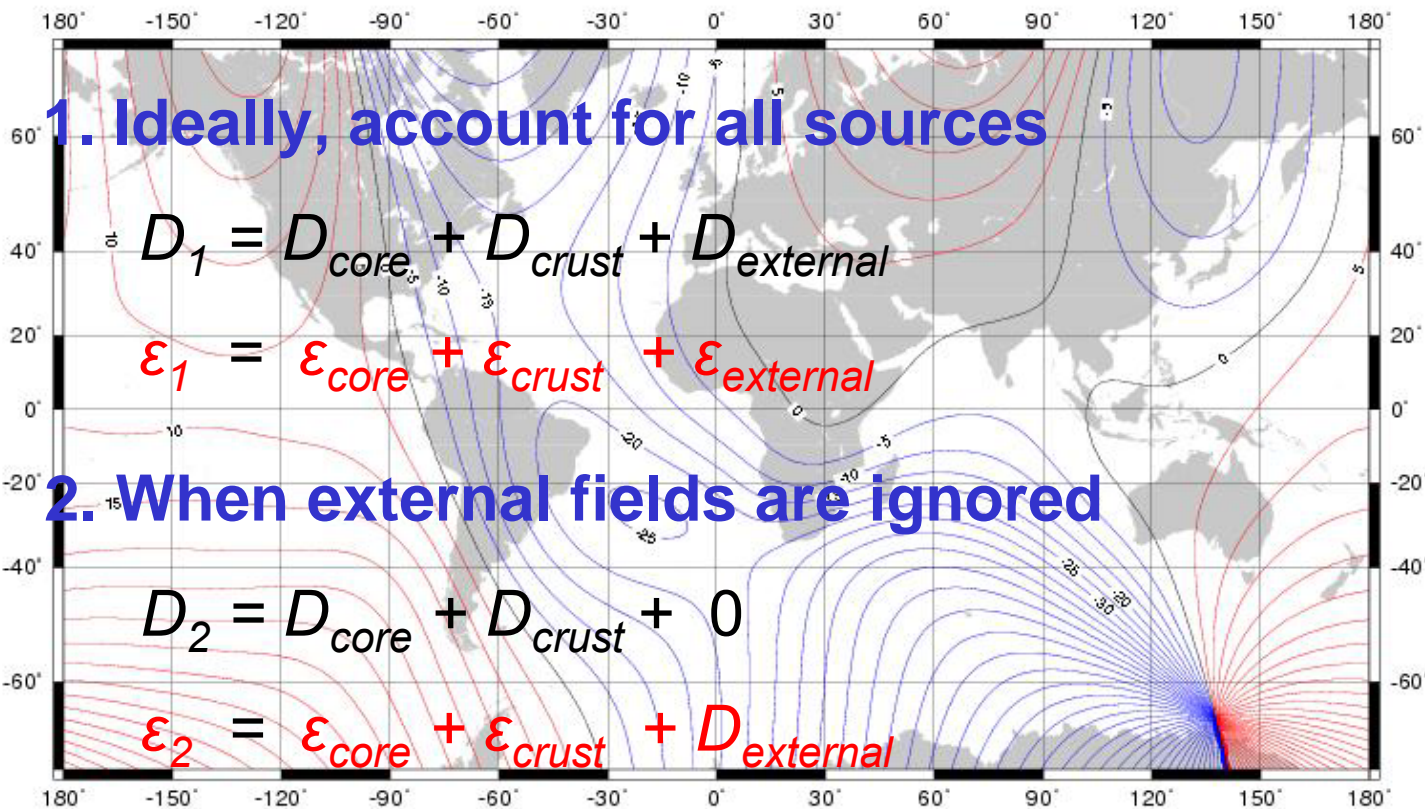
Confidence levels

- Error distributions are not usually normal



- Should not use multiples of σ and assume same confidence as with a normal distribution
- Confidence levels relevant for any error distribution
- Uncertainties presented as limits for confidence levels...
 - 68.3% (equivalent to 1σ if normal)
 - 95.4% (equivalent to 2σ if normal)
 - 99.7% (equivalent to 3σ if normal)

Directional reference and uncertainty



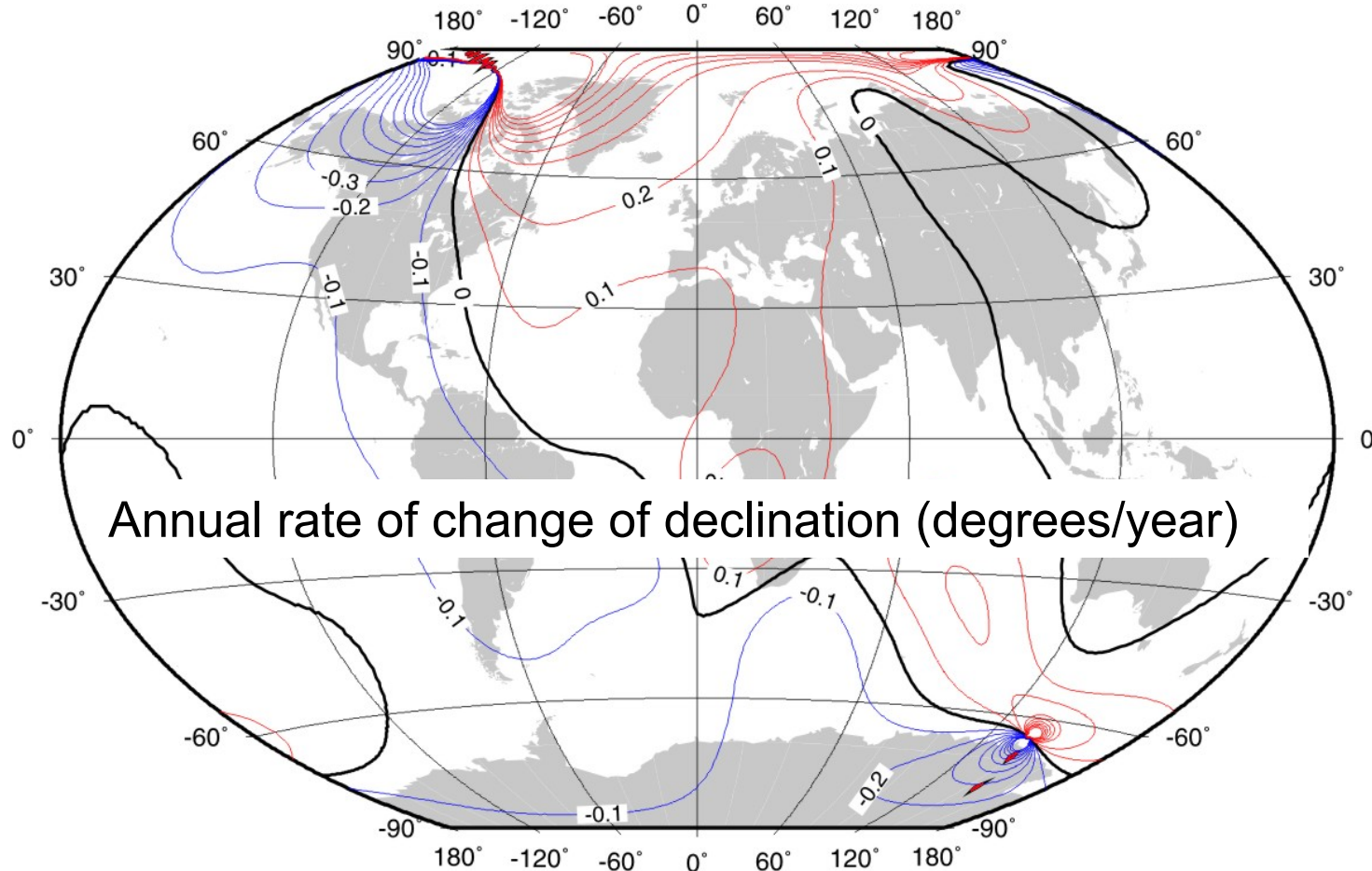
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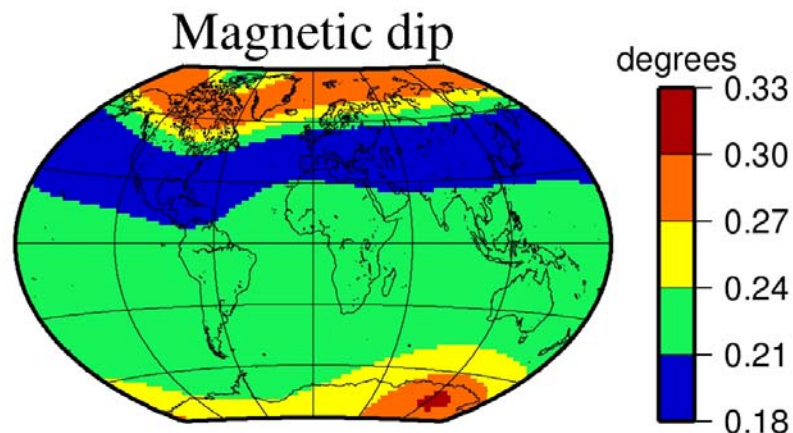
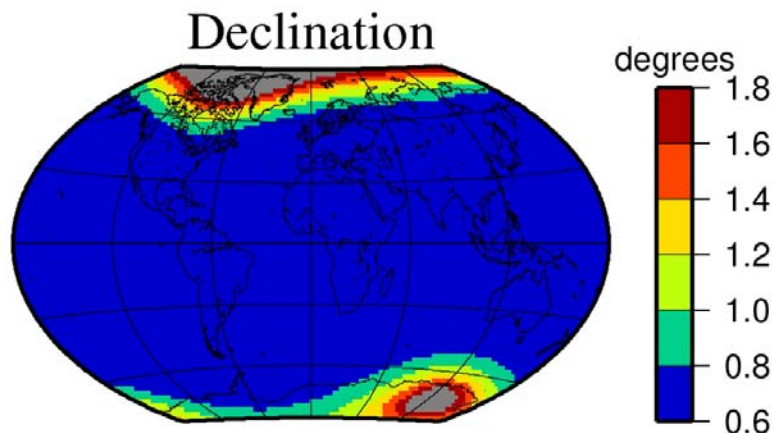
Global magnetic field models

- Spherical harmonic model of the core field and the long wavelength crustal field
- Uses selected satellite and ground based observatory data
- Prediction into future



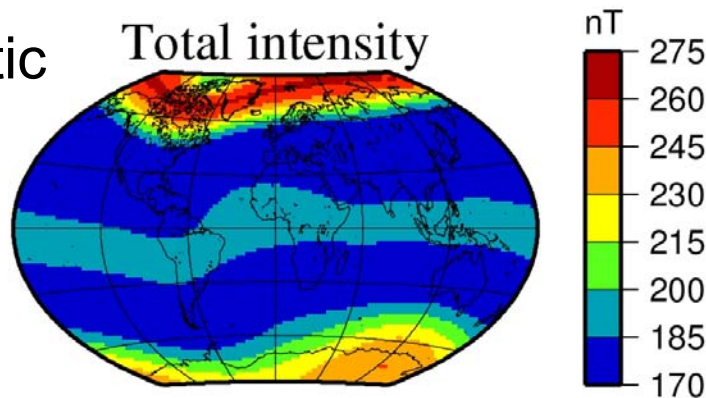
Global field model uncertainties

95.4% confidence limits



Varies with geomagnetic

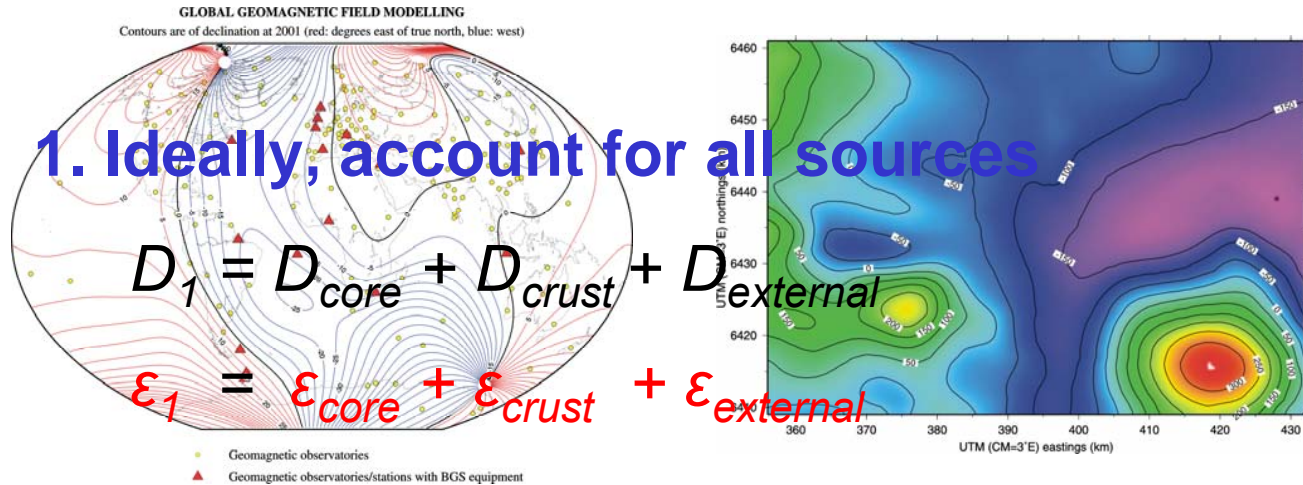
- latitude; and
- activity levels



From Macmillan and
Grindrod, 2010 (SPE
paper 119851)

$$\epsilon_3 = \epsilon_{\text{core}} + D_{\text{crust}} + D_{\text{external}}$$

Directional reference and uncertainty



2. When external fields are ignored

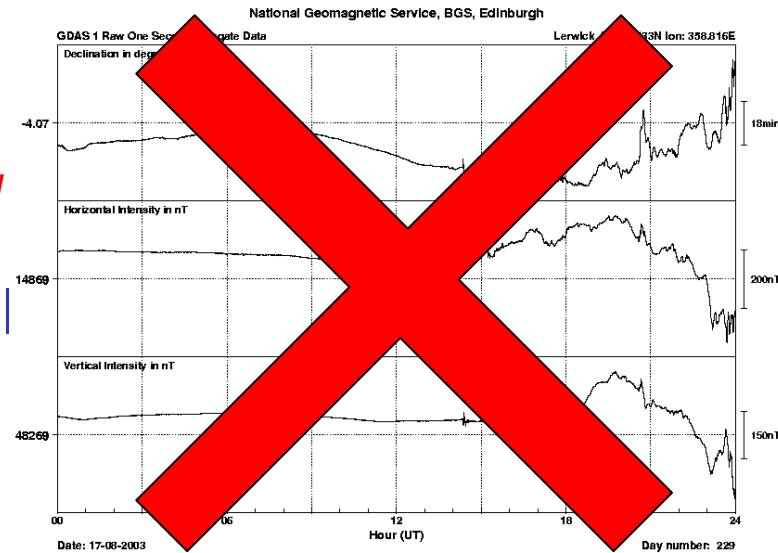
$$D_2 = D_{core} + D_{crust} + 0$$

$$\epsilon_2 = \epsilon_{core} + \epsilon_{crust} + D_{external}$$

3. When crustal and external

$$D_3 = D_{core} + 0 + 0$$

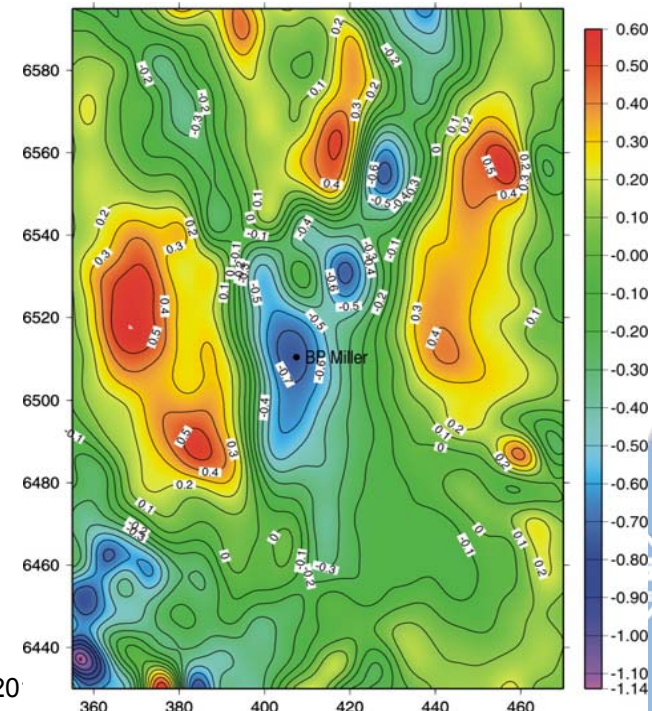
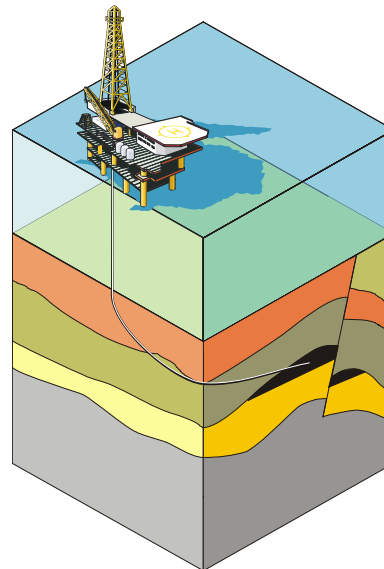
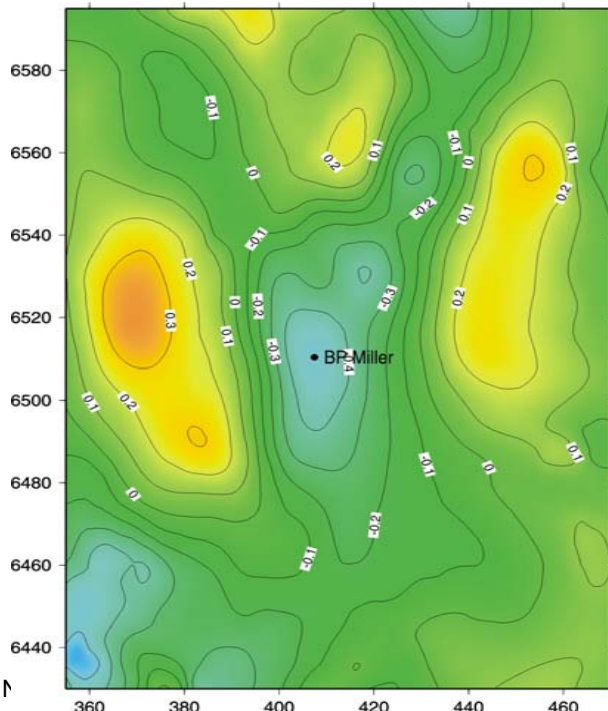
$$\epsilon_3 = \epsilon_{core} + D_{crust} + D_{external}$$



Methods to determine the crustal field

There are three main methods

- Direct measurements of the vector field
- Physical modelling of the magnetic sources
- Transformation of scalar data that exploits the physical properties of magnetic data at or near Earth's surface



Direct measurements

Instruments:

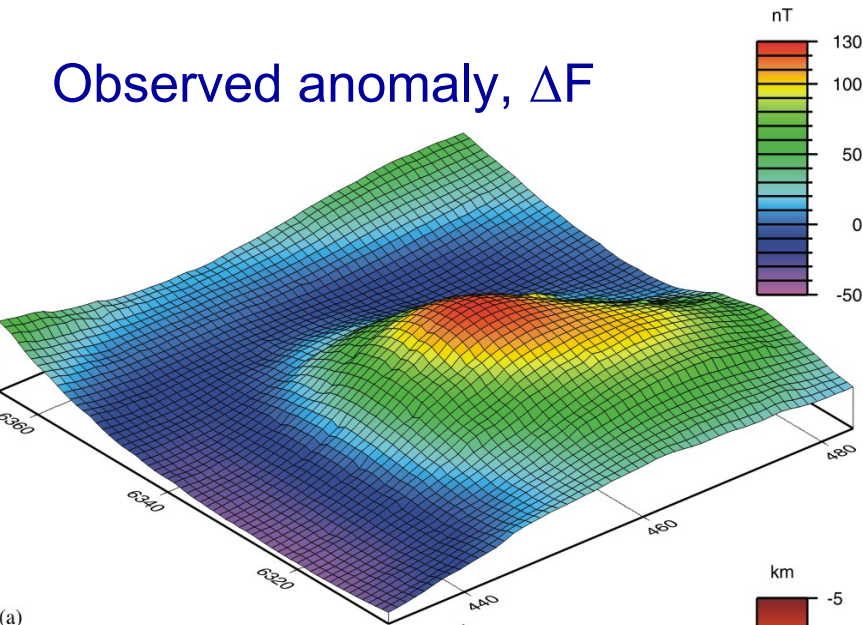
Vector and scalar magnetometers

Ring-laser gyro and GPS

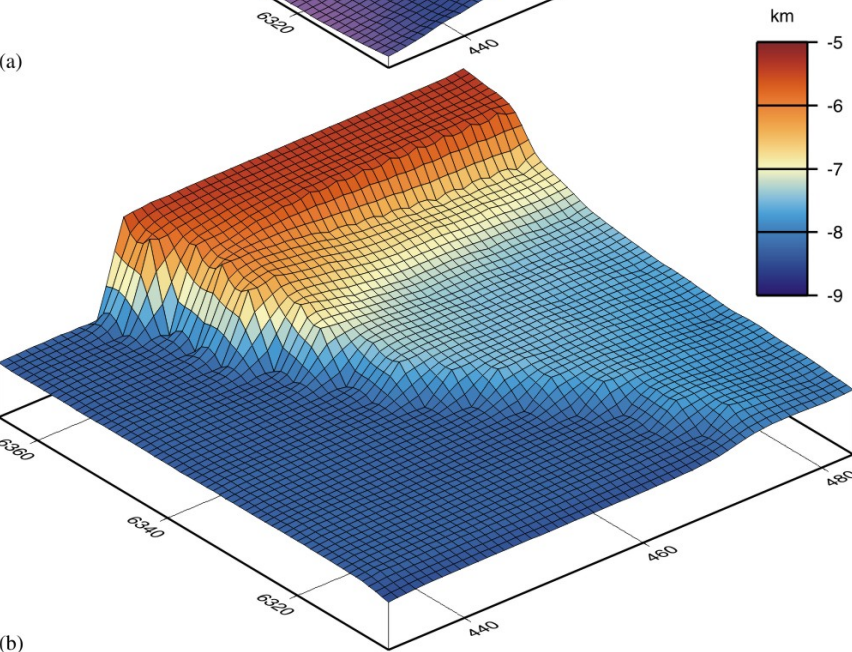


Solving for sources

Observed anomaly, ΔF



(a)

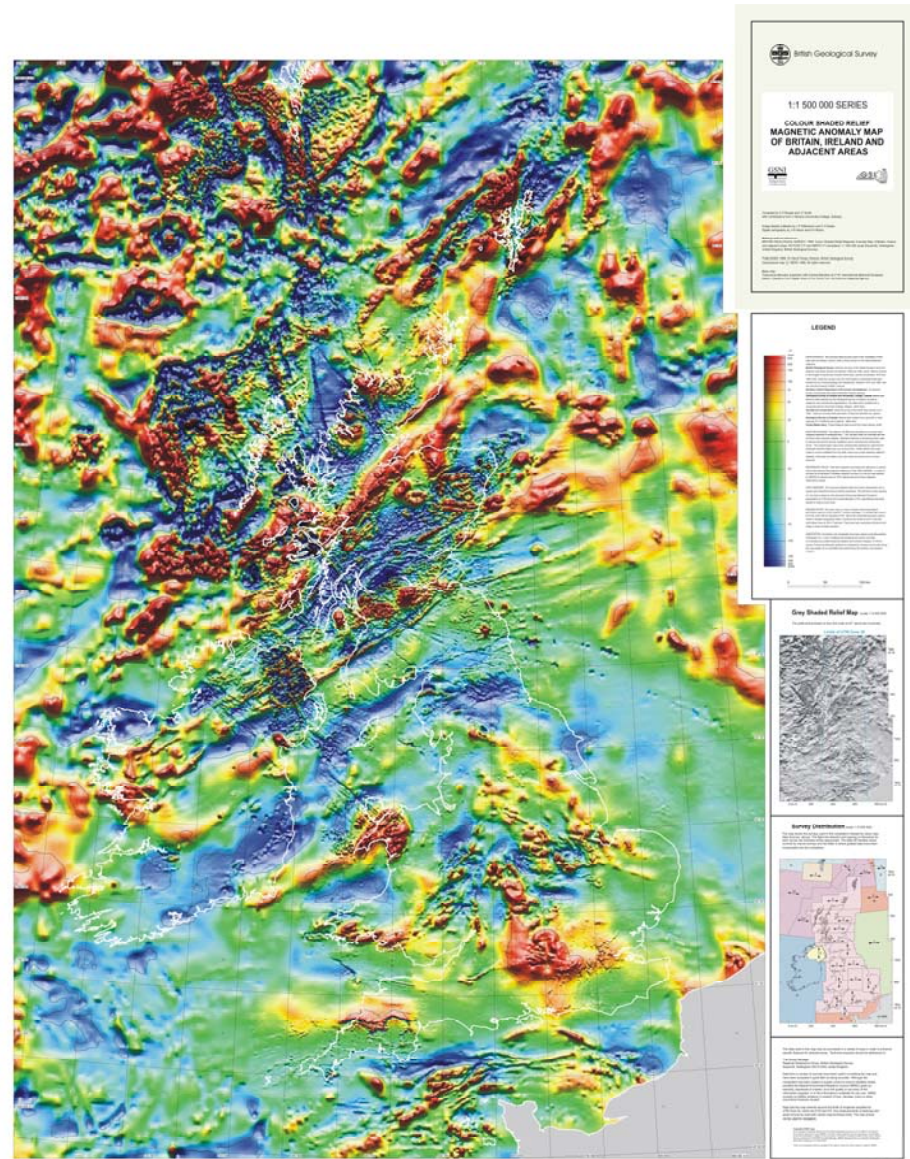


(b)

- Assume magnetisation induced by main field
- Assume magnetisation does not vary with depth

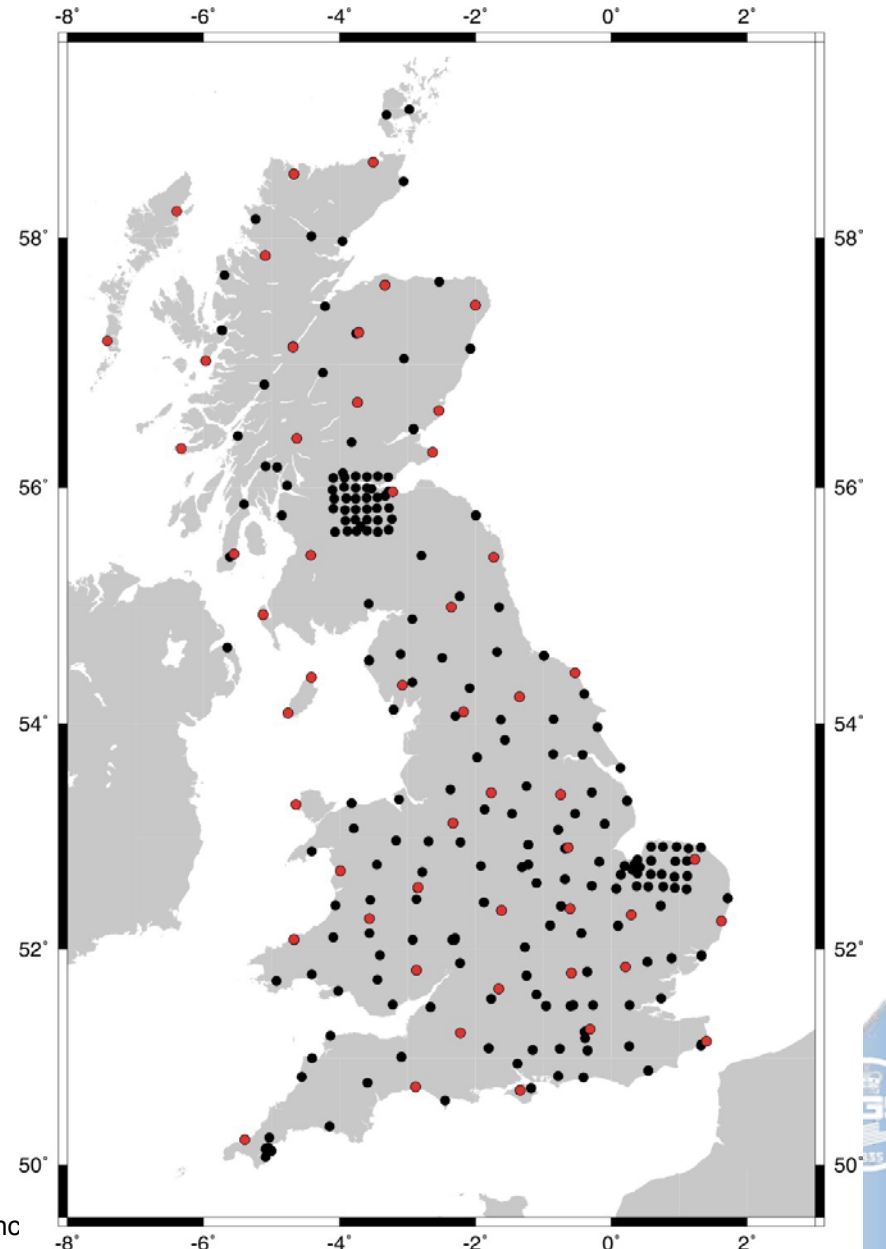
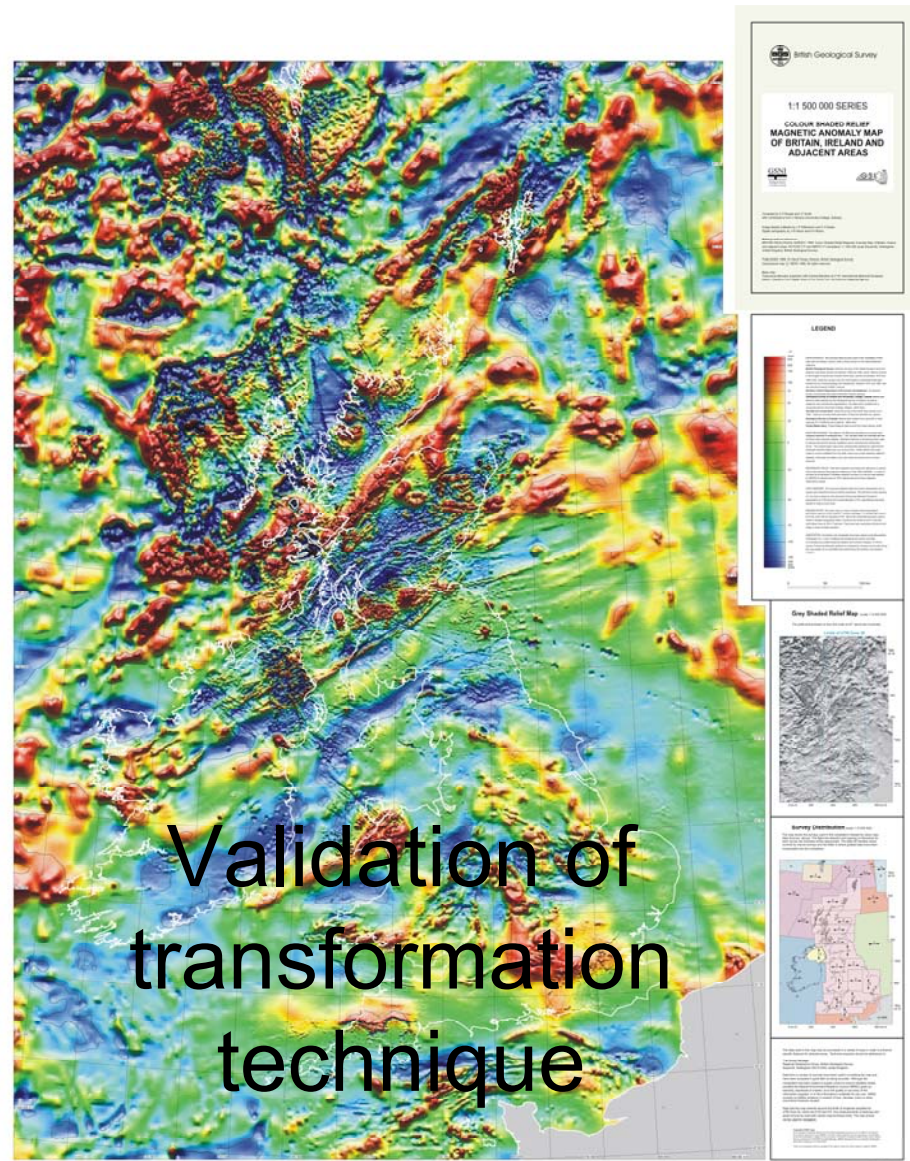
Seismically-determined depth
to magnetic basement

Vector from scalar

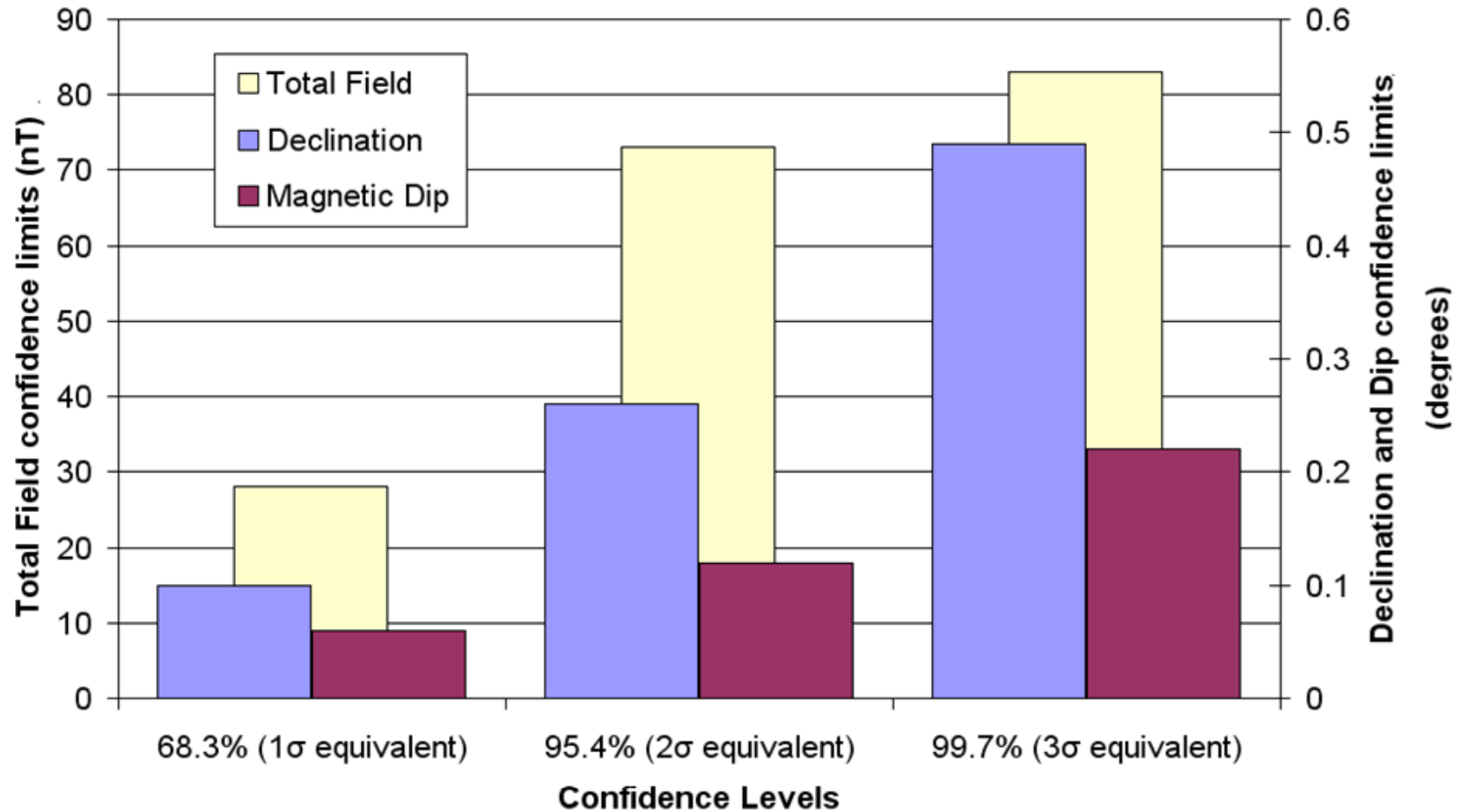


- Assume data collected in source-free region
- Assume constant main field over the area

Crustal (and core) field uncertainties



Crustal (and core) field uncertainties



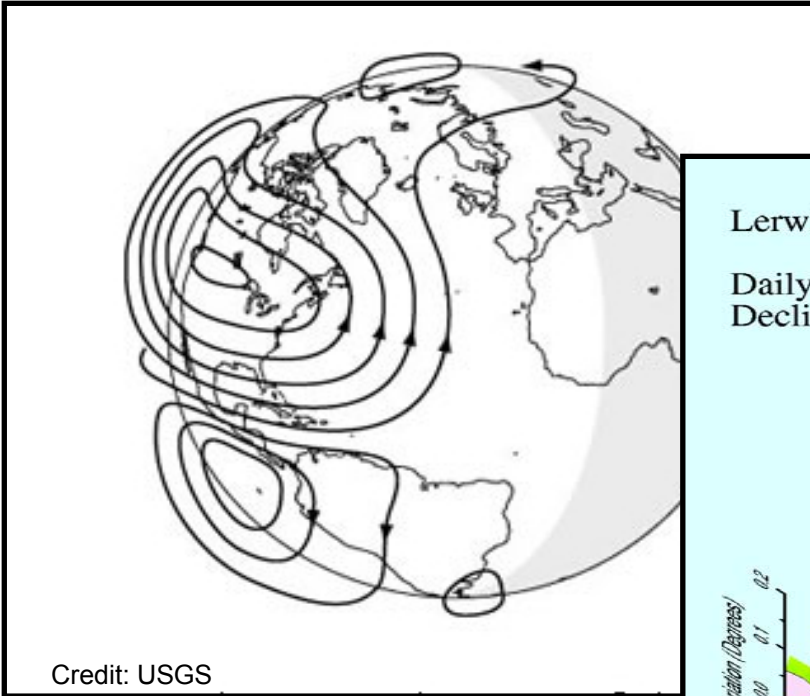
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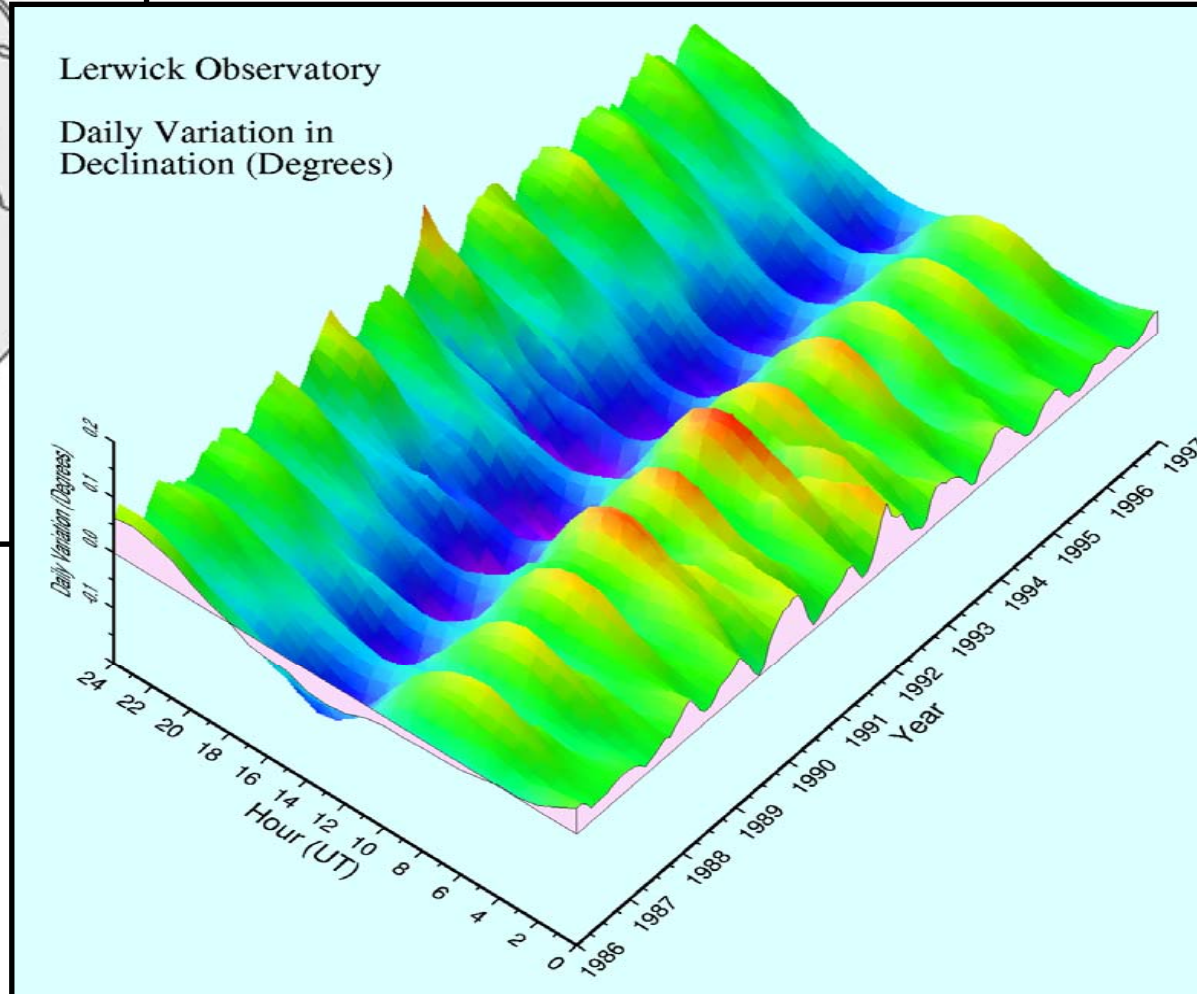
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External fields: 'regular' disturbances

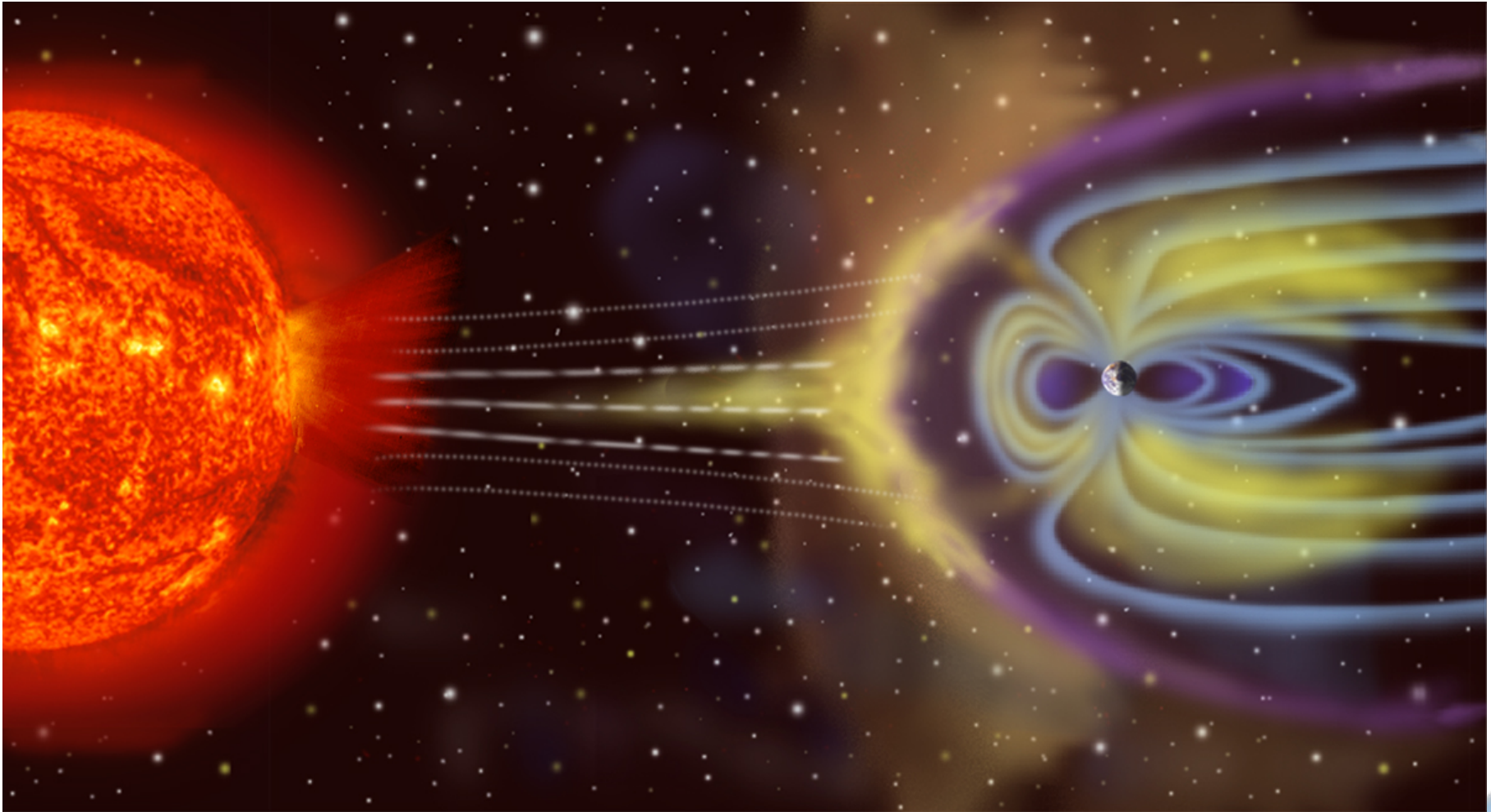


ionospheric currents

diurnal variation



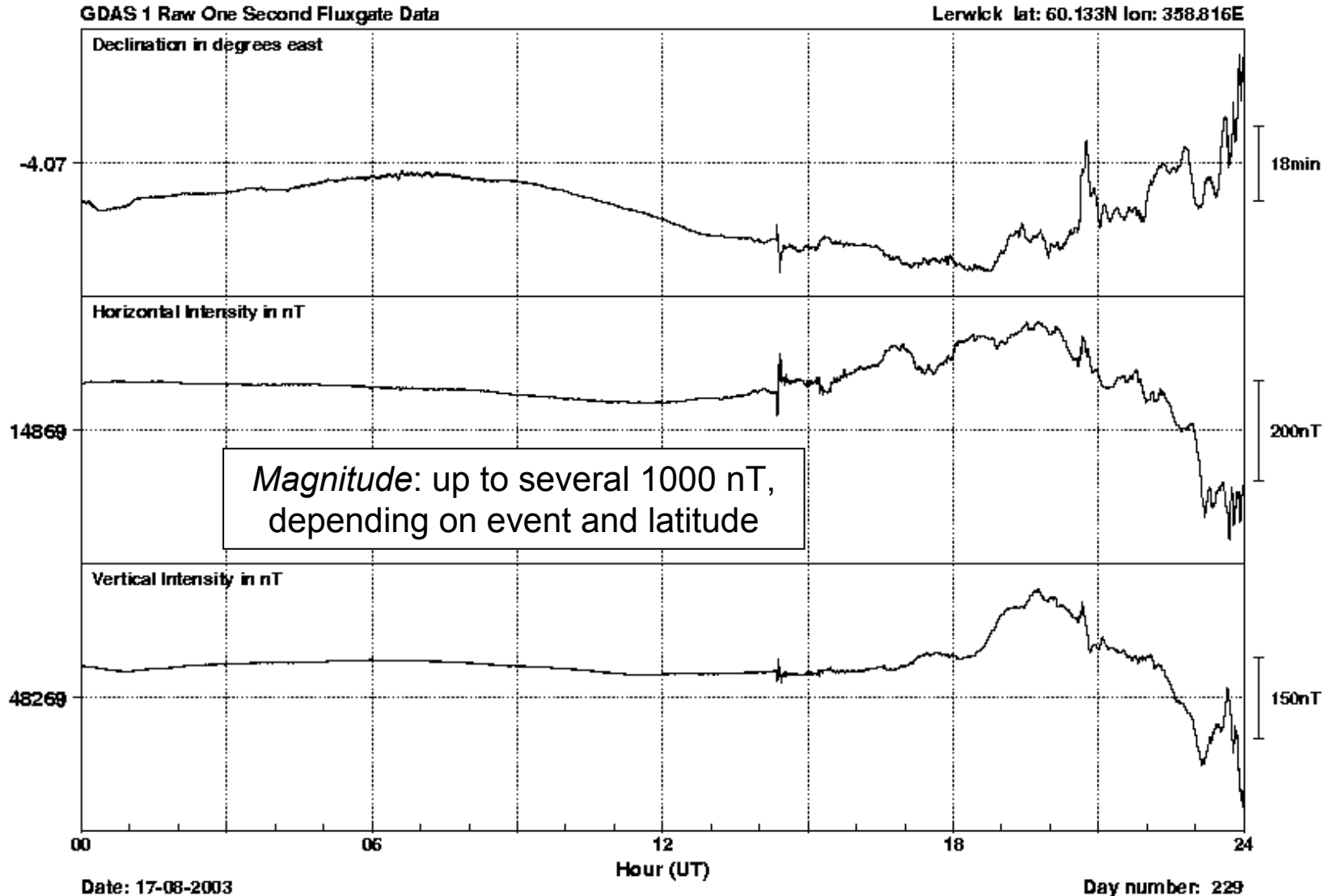
External fields: irregular disturbances



**Solar wind speed 300 to 1000 km/s
(1 - 4 days to travel to the Earth)**

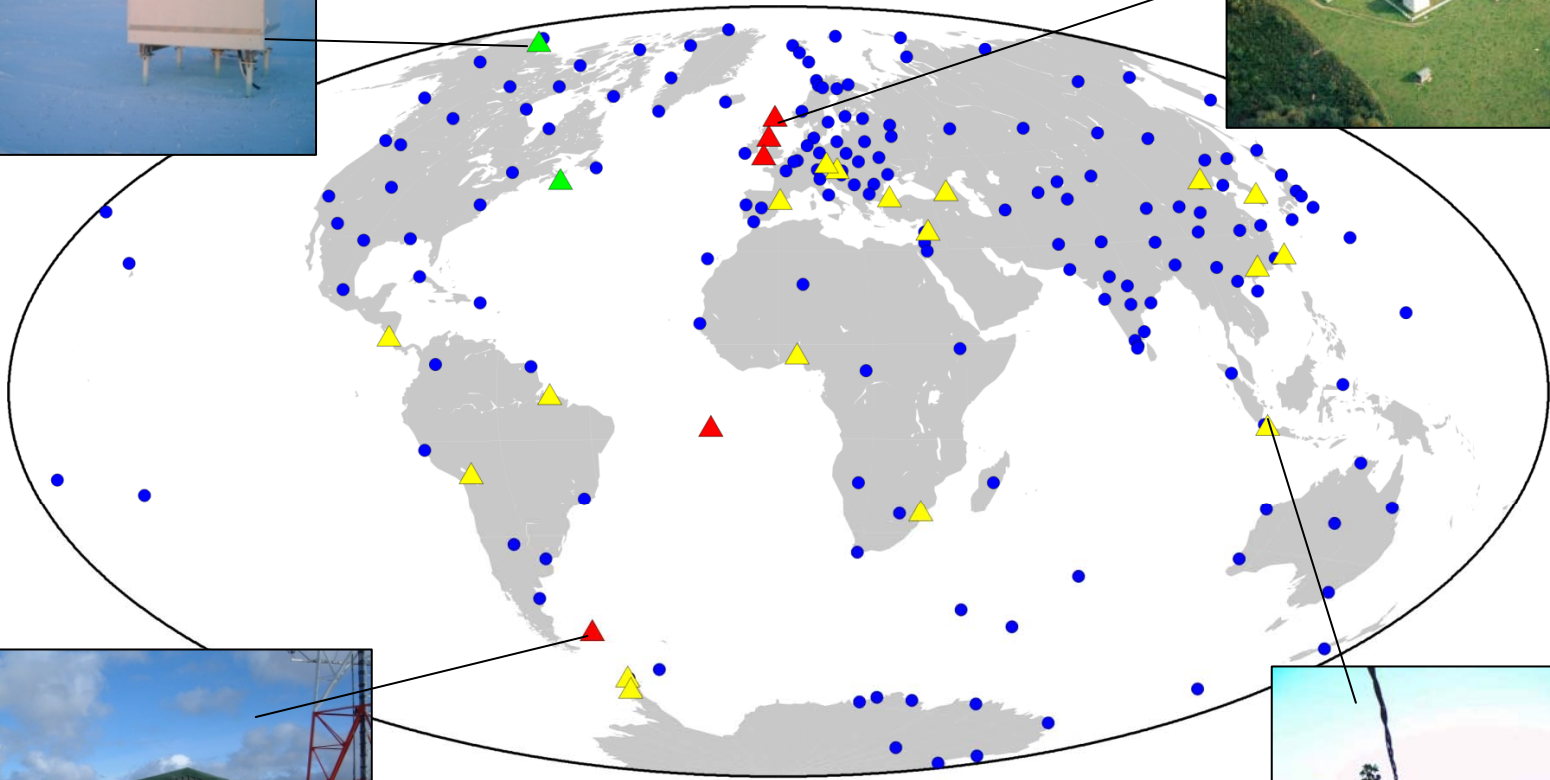
External fields: irregular disturbances

National Geomagnetic Service, BGS, Edinburgh



1/5 years

Geomagnetic Observatories



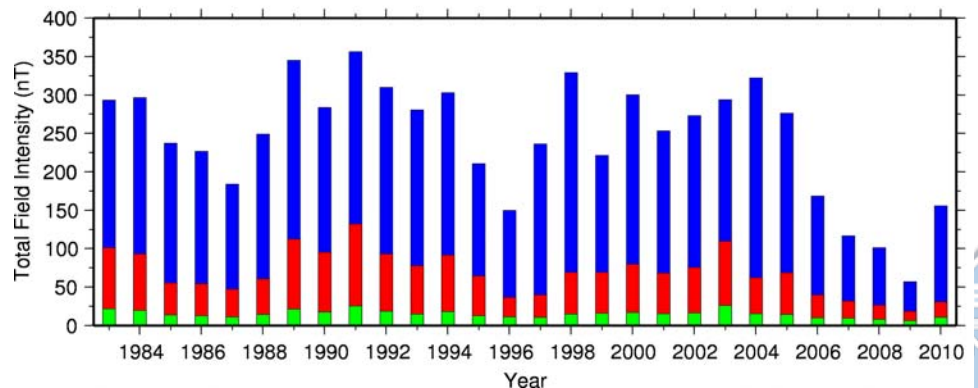
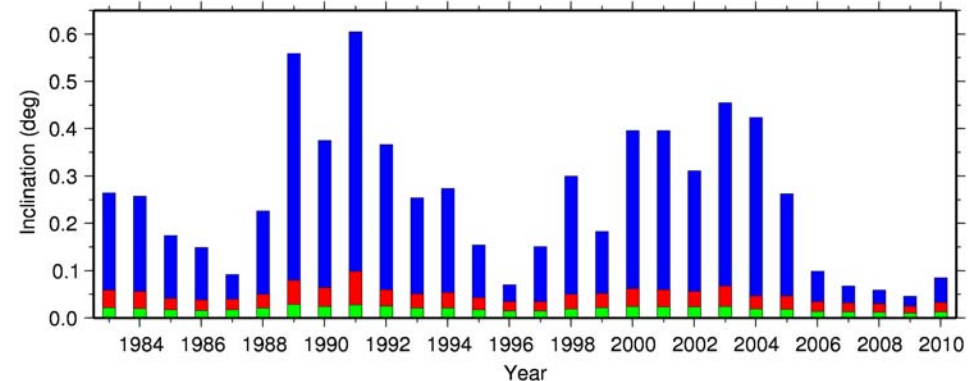
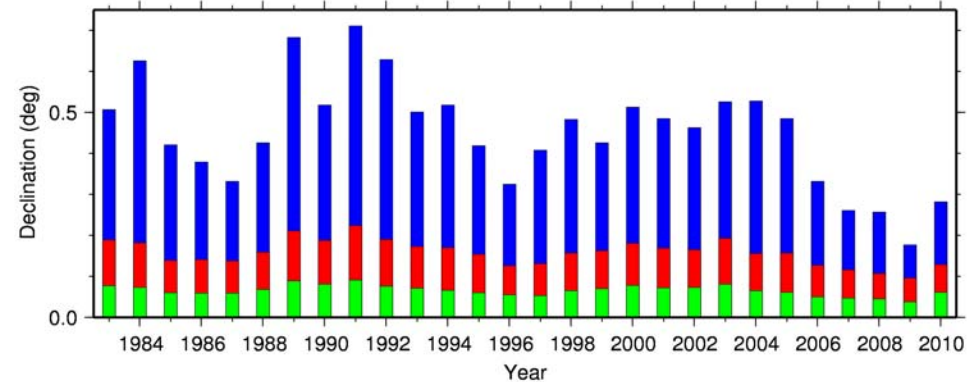
- Geomagnetic observatories
- ▲ BGS observatories
- ▲ BGS/Halliburton observatories
- ▲ Geomagnetic observatories/stations with BGS equip

Crustal (and core) field uncertainties

3-sigma equivalent (99.7%) 2-sigma equivalent (95.4%) 1-sigma equivalent (68.3%)

... due to signal from external fields

variations by year
(at Lerwick)

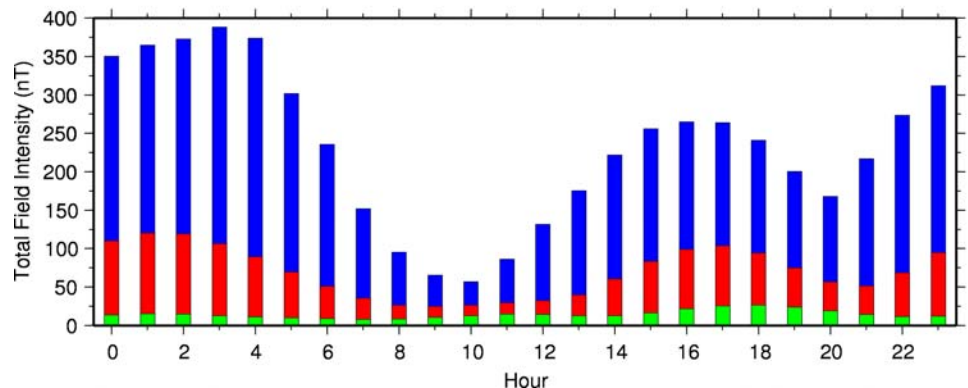
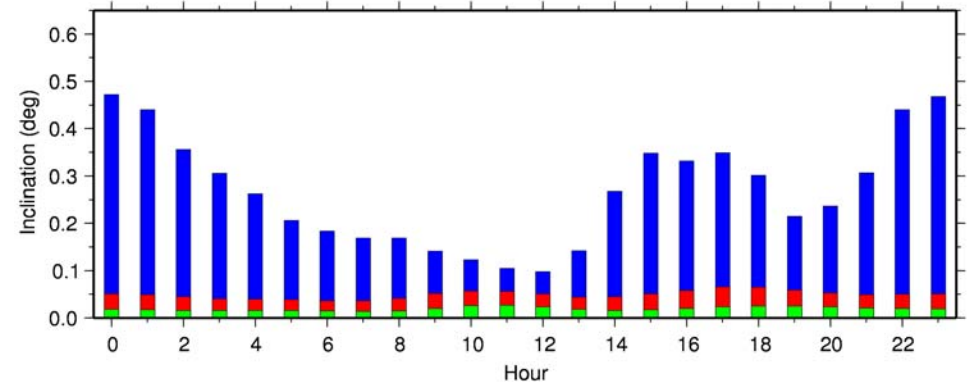
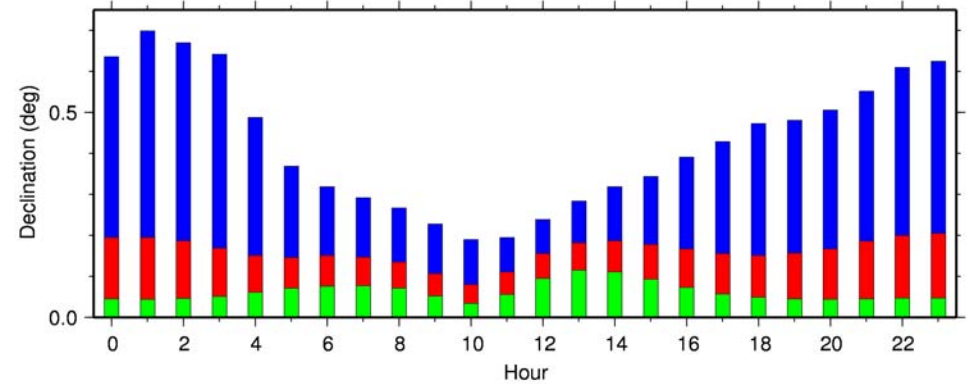


Crustal (and core) field uncertainties

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... due to signal from external fields

variations by hour
(at Lerwick)

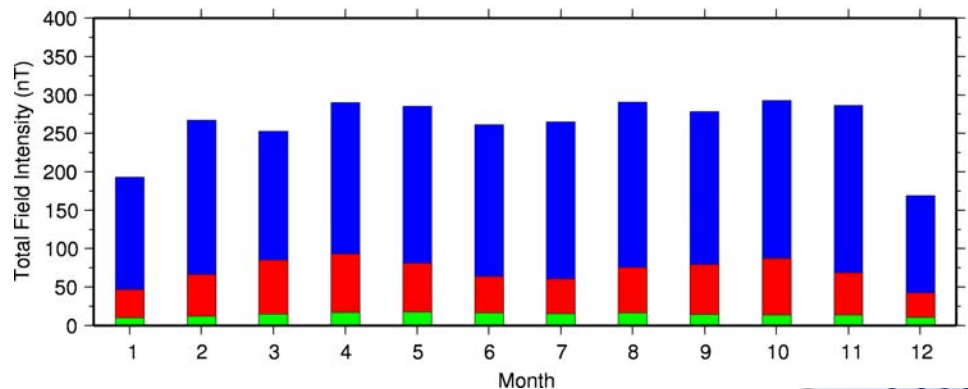
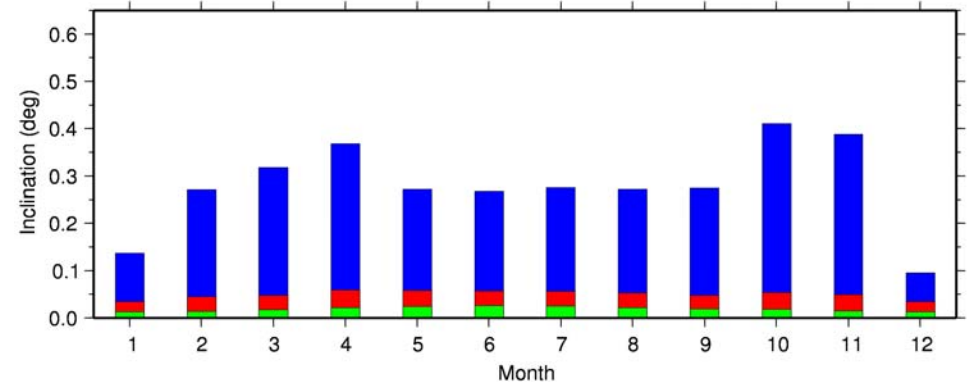
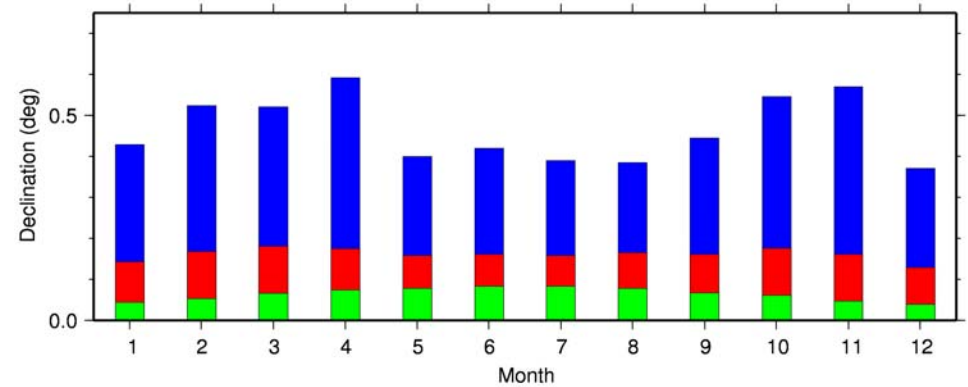


Crustal (and core) field uncertainties

3-sigma equivalent (99.7%) 2-sigma equivalent (95.4%) 1-sigma equivalent (68.3%)

... due to signal from external fields

variations by month
(at Lerwick)



Directional reference and uncertainty

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$$\epsilon_1 = \epsilon_{core} + \epsilon_{crust} + \epsilon_{external}$$

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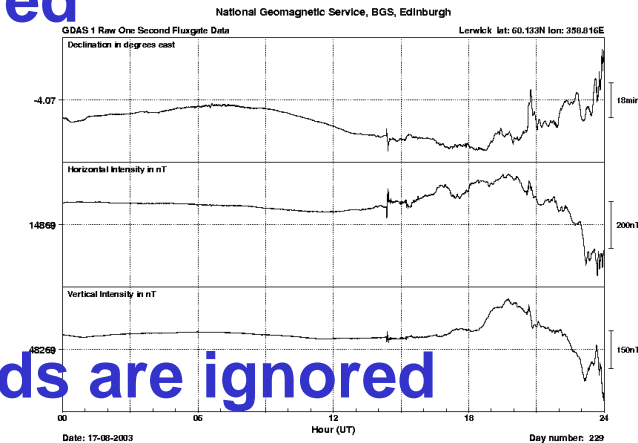
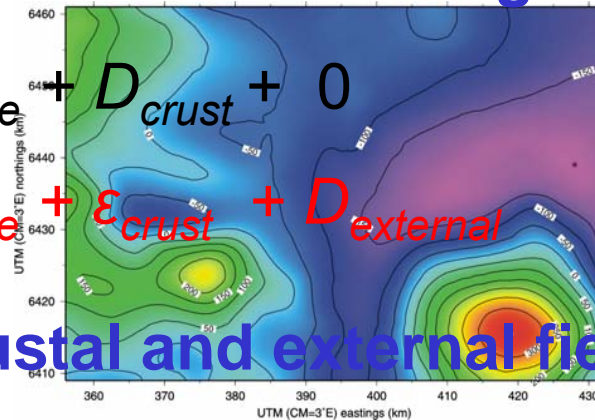
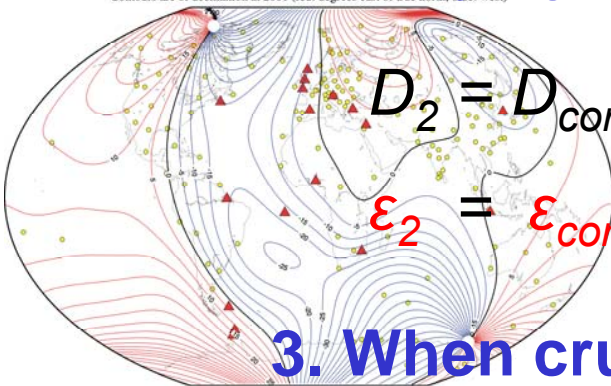
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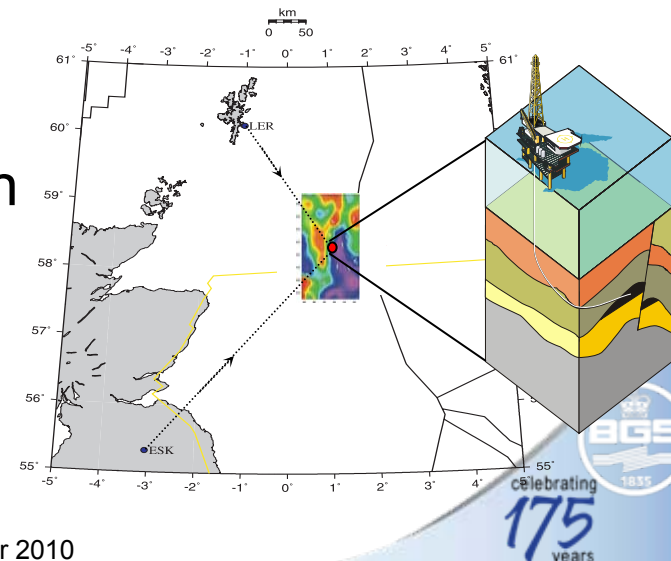
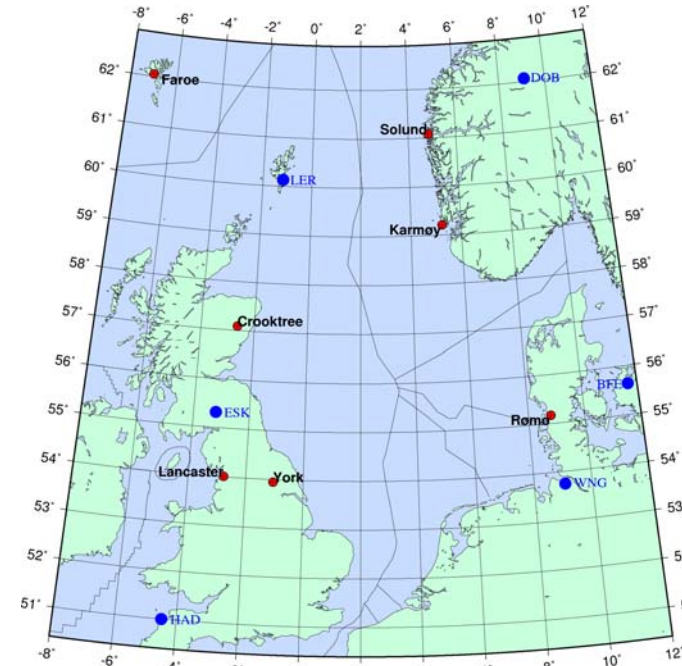
GLOBAL GEOMAGNETIC FIELD MODEL IGRF11
Contours are of declination at 200 nT intervals, east or west



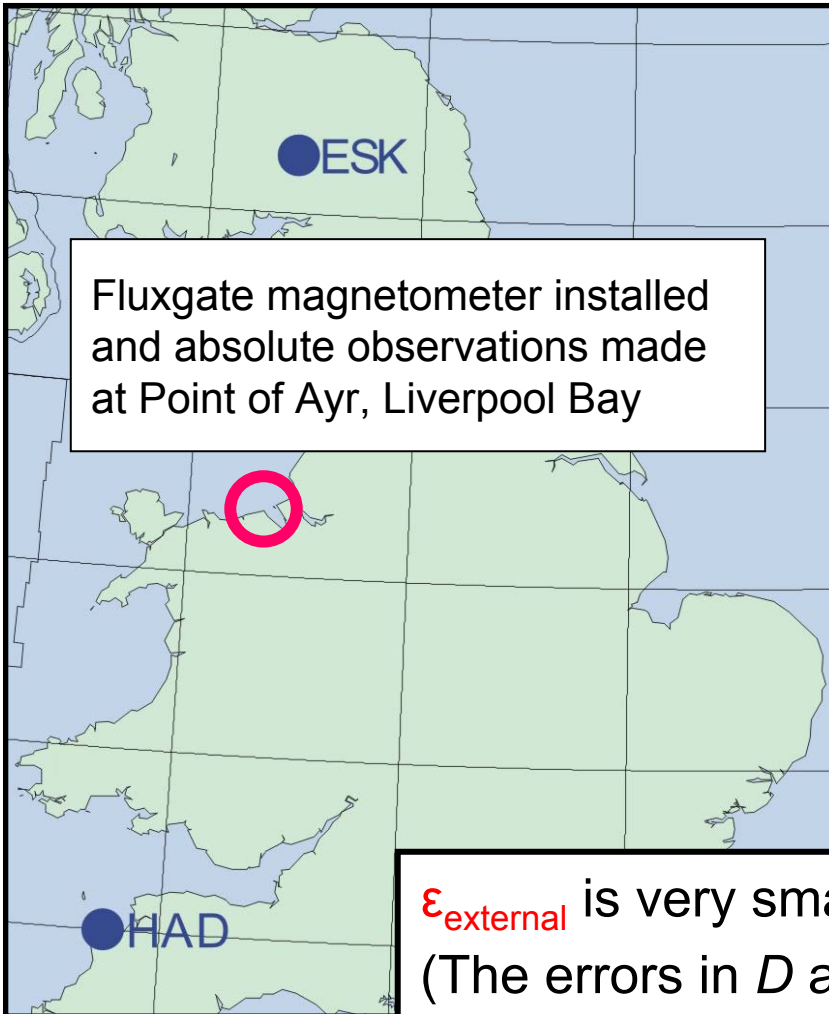
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Estimating the external field at the drill site

- Use data from nearby magnetic observatories (and/or calibrated variometer stations)
 - can use one or more
- Take advantage of observatory quality control, quasi-definitive data and real-time operations
- Use the real observatories to create a 'virtual' observatory at the drilling location



Liverpool Bay experiment 1994-95



Month	D (deg)	I (deg)	F (nT)
Oct 94	0.006	0.005	5.3
Nov 94	0.008	0.006	3.4
Dec94	0.007	0.003	5.3
Jan95	0.005	0.002	4.4

Russell, Shiells and Kerridge, (1995)

$\epsilon_{\text{external}}$ is very small compared with the errors $\epsilon_{\text{core}} + \epsilon_{\text{crust}}$
(The errors in D and I are $\sim 0.07^\circ$ and $\sim 30\text{nT}$ in F .)

At North Sea latitudes the contribution of $\epsilon_{\text{external}}$ does not significantly affect the overall uncertainty of the technique.

External field uncertainty revisited 2005

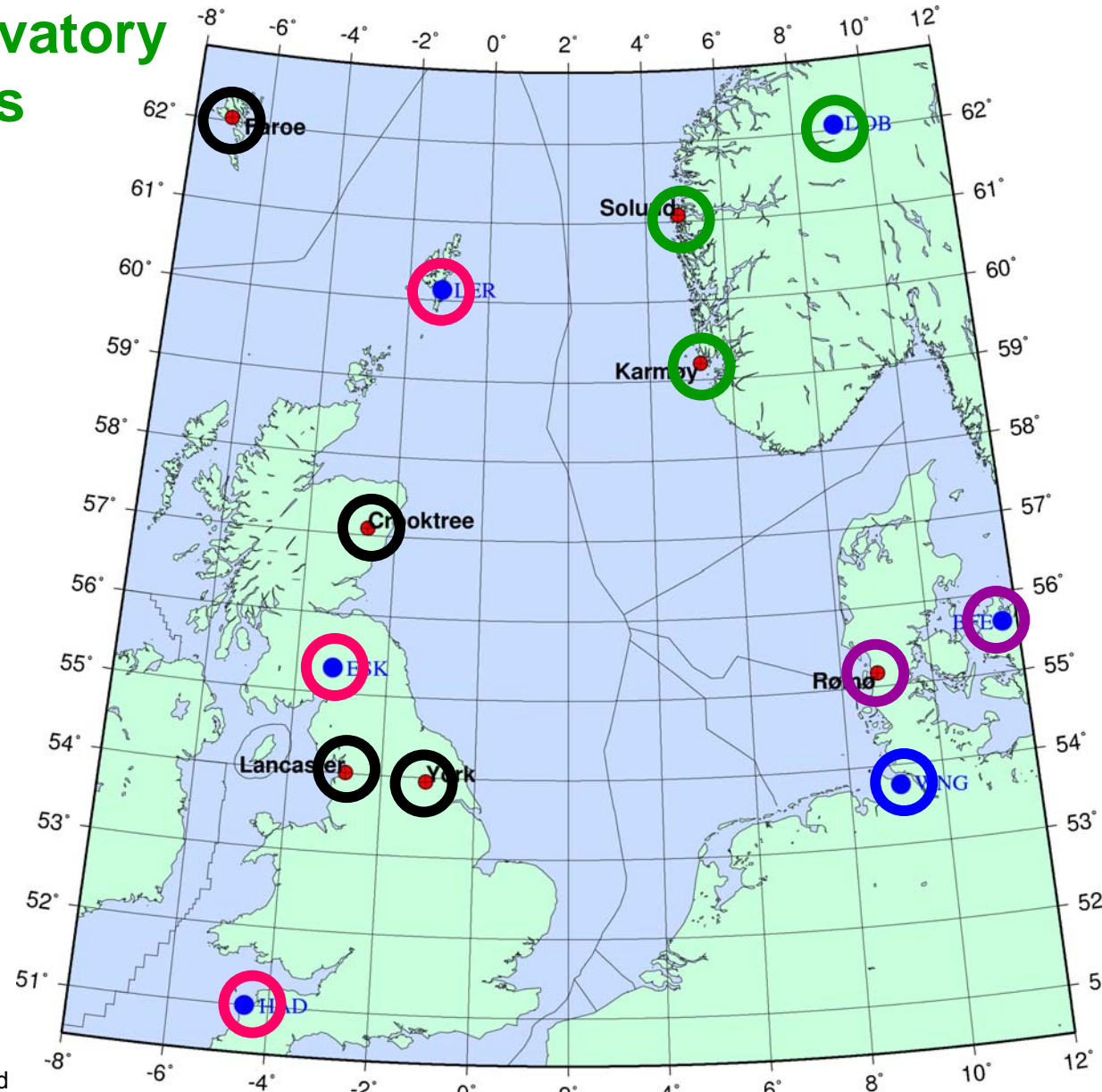
**TGO (Norway) observatory
& variometer stations**

**GFZ (Germany)
observatory**

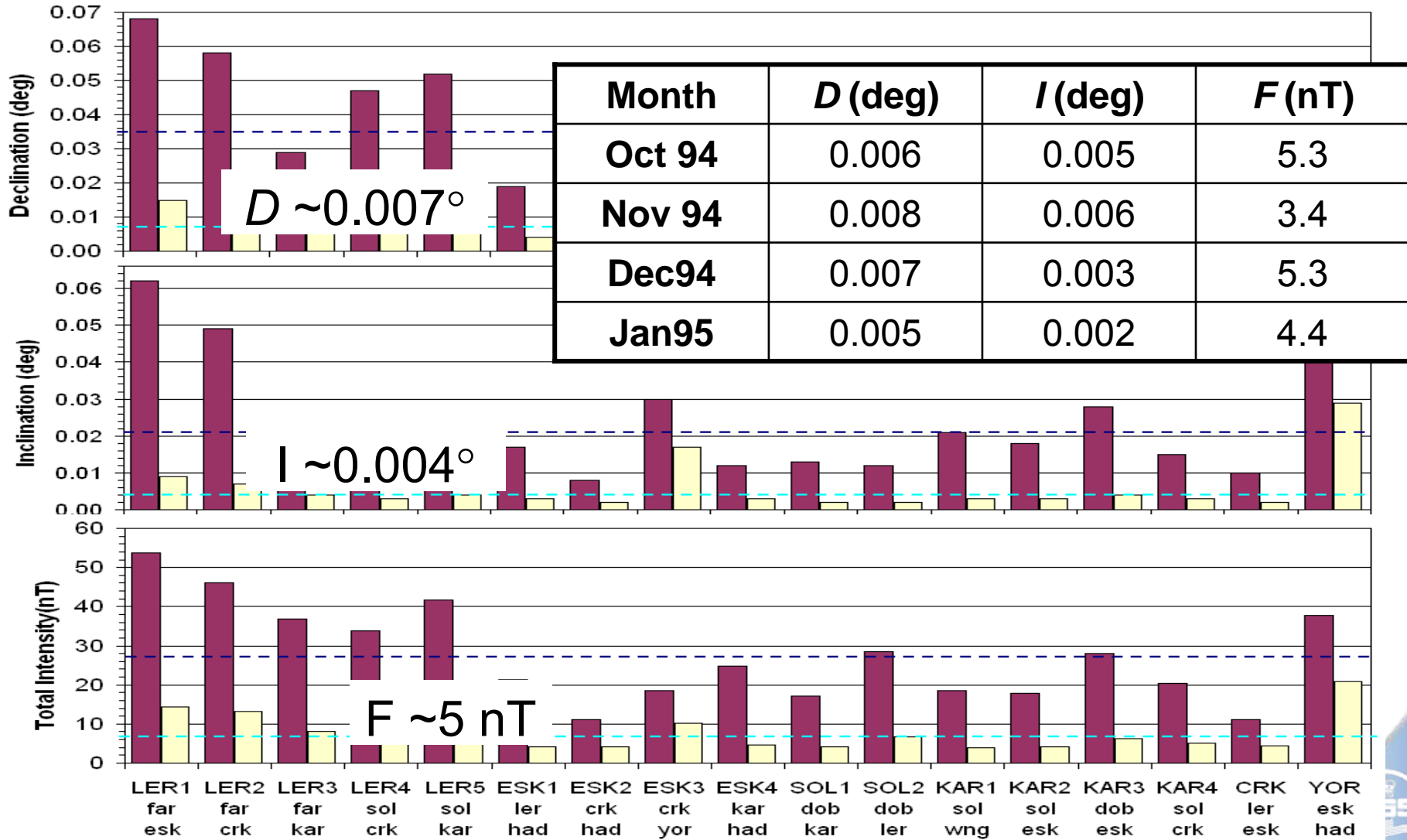
**DMI (Denmark)
observatory &
variometer station**

**SAMNET (UK)
variometer stations**

**BGS (UK)
observatories**



External field uncertainties ($\epsilon_{\text{external}}$) around the North Sea

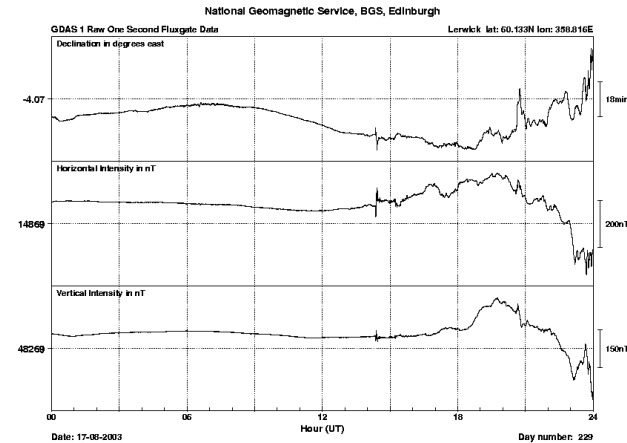
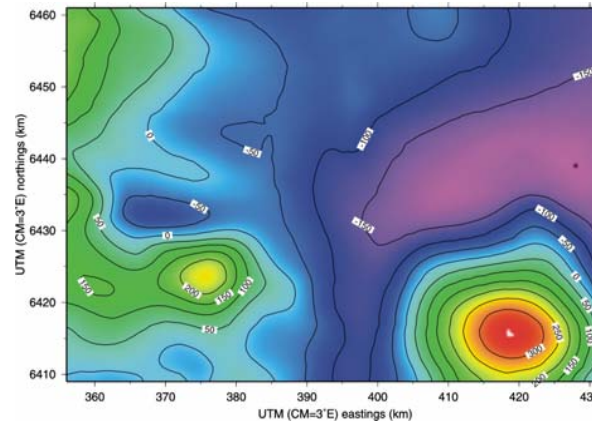
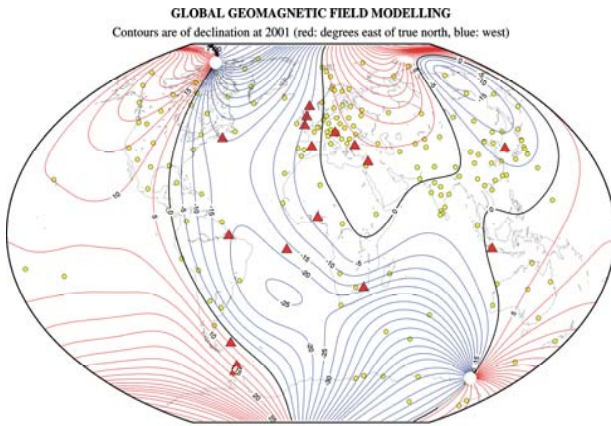


Directional reference and uncertainty

1. Ideally, account for all sources

$$D_1 = D_{core} + D_{crust} + D_{external}$$

$$\epsilon_1 = \epsilon_{core} + \epsilon_{crust} + \epsilon_{external}$$



Summary and Conclusions

- Providing a single answer on the uncertainty associated with the geomagnetic reference field is hard.
- The uncertainty associated with a particular reference value, depends on
 - what sources are included in the estimate;
 - location on the earth's surface; and often on
 - time of day, season and solar cycle
- Including all sources will
 - reduce the uncertainties; and importantly,
 - provide a more robust estimate of the uncertainties by removing any hourly, seasonal and solar cycle variations

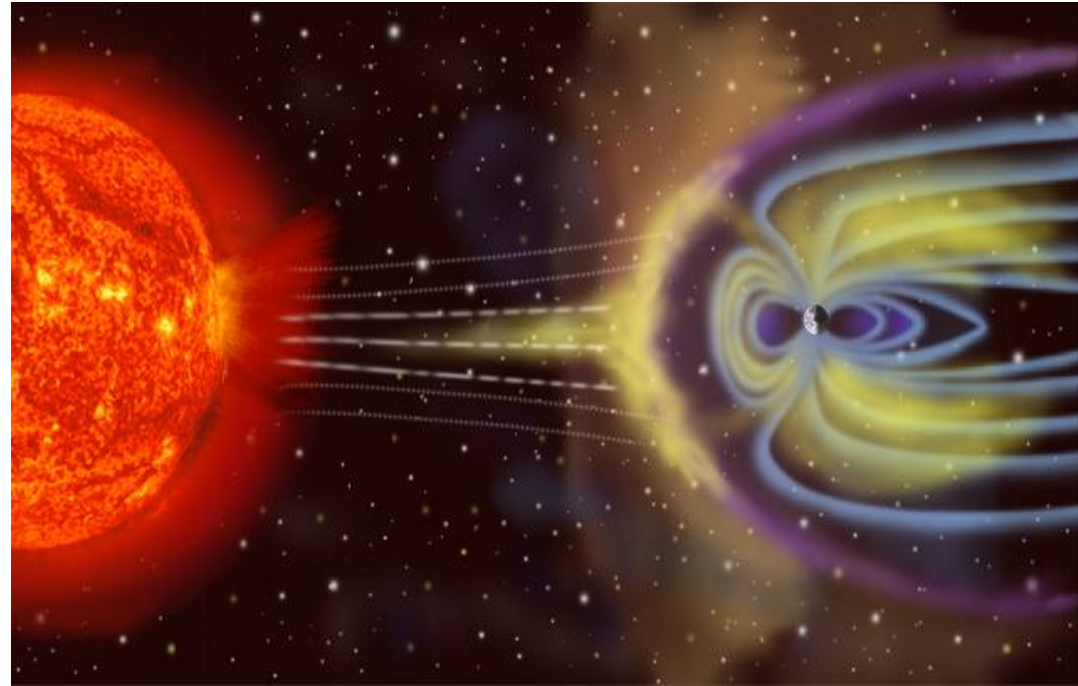


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External field sources

- Ionospheric currents
 - from about 100 km altitude
- Magnetospheric currents
 - out to several Earth radii
- Secondary currents induced in the earth



External fields: long-term trends

Sunspot Cycle and Annual Number of Magnetic Storms

