Managing Main & Crustal Magnetic Fields and New Developments in Global Magnetic Modeling

- Wellbore position uncertainty
- SD and HD geomagnetic models
- In-Field Referencing methods
 - Shortcomings of plane grid methods
 - Global ellipsoidal harmonic method



Texas and Alberta aeromagnetic survey examples

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Positional Uncertainty in MWD

- Largest source of lateral error: Magnetic field
- Accurate geomagnetic models and advanced corrections significantly reduce this error



Ellipses of Uncertainty Study



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Uncertainties in Middle East

Lateral uncertainty at TD for well with 11,000 ft lateral section

Well Azim.	MWD+IGRF	MWD	MWD+HD	MWD +IFR1	MWD +IFR1+MS
	(ft)	(ft)	(ft)	(ft)	(ft)
E	370 (+7%)	345	325 (-6%)	289 (-16%)	145 (-58%)*
SE	329 (+10%)	300	277 (-8%)	233 (-22%)	134 (-55%)
S	266 (+16%)	229	199 (-13%)	131 (-43%)	117 (-49%)

*with limitations

Vertical Uncertainty at TD

TVD	MWD	MWD+IFR1+SAG+MS
10600 ft	119 ft	71 ft (-40%)

3D Ellipsoids given for 95% confidence interval = 2.79 sigma

Error model: ISCWSA OWSG 2014

Main Field Commission Error

Only shown is the commission error of the main field. This does not include errors from omitting the crustal field.



This demonstrates the importance of using annually updated models

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Crustal Anomalies: The Old View



Geomagnetic Power Spectrum



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Representing the magnetic field

• Magnetic field vector represented as gradient of a potential

 $\boldsymbol{B} = -\boldsymbol{\nabla} V$

• Potential satisfies LaPlace's differential equation:

 $\nabla^2 V = 0$

- Harmonic functions satisfy LaPlace's equation
 - Sphere: Spherical harmonics
 - Ellipsoid: Ellipsoidal harmonics
 - Plane: Harmonic waves

Spherical harmonic Models

$$\mathbf{B} = -\nabla V$$

Spherical harmonic expansion of potential V

 $V(\lambda,\psi,r,t) = a \sum_{n=1}^{N} \sum_{m=0}^{n} \left(\frac{a}{r}\right)^{n+1} \left(g_n^m(t)\cos m\lambda + h_n^m(t)\sin m\lambda\right) P_n^m(\sin\psi)$

N = *Degree of the model*

a = Geomagnetic reference radius (6371.2 km)

 $g_n^m(t)$ and $h_n^m(t)$: Model coefficients

 $P_n^m(\sin\psi)$: Associated Legendre functions

Standard Definition Models (MWD)



*Degree of BGGM depends on evaluation date

Declination Comparison (Bakken)

MVSD

Main Field



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SD models: Spherical harmonic expansion of potential

$$V(\lambda,\psi,r) = a \sum_{n=1}^{N} \sum_{m=0}^{n} \left(\frac{a}{r}\right)^{n+1} \left(g_n^m \cos m\lambda + h_n^m \sin m\lambda\right) P_n^m(\sin\psi)$$

High Definition models: Ellipsoidal harmonic expansion

$$V(\lambda,\beta,u) = a \sum_{n=1}^{N} \sum_{m=0}^{n} \frac{Q_n^m(i\frac{u}{E})}{Q_n^m(i\frac{b}{E})} \left(g_n^m \cos m\lambda + h_n^m \sin m\lambda\right) P_n^m(\sin\beta)$$

- u = semi-minor axis of confocal ellipsoid at location
- *b* = semi-minor axis of WGS84 ellipsoid
- E = distance of foci from Earth center
- β = reduced latitude
- Q_{nm} = Associated Legendre functions of the second kind

High Definition Models (MWD+HD)^{Slide 13 of 34}



Declination Comparison (Bakken)



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In-Field Referencing (MWD+IFR)



IFR models combine local marine or airborne magnetic total field measurements with global satellite magnetic data

Solving for the IFR Model

- Magnetic field vector $B = -\nabla V$
- Aeromag surveys measure the total field F, which is the derivative of the potential in the direction of the magnetic field:

 $\mathbf{F} = -\boldsymbol{\nabla} V \cdot \boldsymbol{b}$

• For a local IFR model, this is not enough information to solve LaPlace's differential equation:

$$\nabla^2 V = 0$$

• For a unique solution, have to specify boundary conditions!

Methods Using a Plane Grid

- FFT method (e.g. Dean 1958; Russel, Shiells and Kerridge, SPE 30452, 1995)
- Equivalent source method (Dampney, 1969; Macmillan & Billingham, ISCWSA-40, 2014)
- Both techniques assume that all anomalies are completely contained within the local grid
 - Equivalent to assuming V = 0 on the boundary



Russel, Shiells and Kerridge, SPE 30452, 1995



Macmillan & Billingham, ISCWSA-40, 2014

Assumption of Plane Grid Methods



- Grid methods assume that all anomalies are confined to the local grid.
- This implies that there are no wavelengths longer than the grid size.

Assumption of Plane Grid Methods



In reality, the crustal field has significant power at long wavelengths At those wavelengths, the plane approximation is no longer valid

Synthetic Example

- Use 80 km x 80 km grid (shown in blue)
- To make it simple, assume that dF = 0 on the grid
- Surrounded by a long-wavelength 400 nT anomaly



Plane Grid Method Error





Mathematical reason for error: Grid method assumes that the potential is zero on the grid boundary

Any method used to solve LaPlace's equation on a grid without specifying boundary conditions will give incorrect results

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Accurate IFR Solution

Problem: Magnetic potential not known on boundary Solution: Eliminate the boundary

• This is a common approach used in weather models



Compute the IFR model as a global model using ellipsoidal harmonics (MagVAR IFR method, e.g. IADC/SPE-150107, 2012)

• Also used to compute the true solution for the previous example

In-Field Referencing (MWD+IFR)



An accurate IFR model accounts for the entire spectrum without gaps

Declination Comparison (Bakken)



Why not just measure the vector at the surface?

Earth surface

9,000 ft TVD





Magnetic anomalies increase with depth

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- Even surface directional measurements need downward transformation
- Still need to solve
 LaPlace's equation

- Surface directional measurements are valuable
- For an accurate reference field at depth one still needs a transformation

Secular Variation Correction: The Old Way Slide 26 of 34



Often a "crustal correction" is added to a main field model for the desired date. This will lead to disagreements

- Between different main field models
- If the degree of the main field model changes

Example: BGGM2004 = degree 20 BGGM2007 = degree 30 BGGM2010 = degree 50 BGGM2012 = degree 40 BGGM2014 = degree 50

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Secular Variation Correction



An accurate method of correcting for change in main field:

- 1. Start with complete field given for reference date
- 2. Subtract a yearly updated model for this reference date
- 3. Add the same model for the desired drilling date

Note: This does not work if the model degree changes with date: Say we have a ground shot from 1998. By subtracting BGGM2014 for 1998 and adding BGGM2014 for 2015 we are in fact subtracting a degree 13 model and adding a degree 50 model, which is invalid.

Permian Basin Aeromag Example







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Alberta Oil Sands Aeromagnetic Survey

- 1000 m spacing
- 250 m spacing
 - Cold Lake Weapons Range



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ore Accurate

Relevant Tool Codes for Well Planning and Anti-Collision Scans

Tool Code	Magnetic reference model	Survey corrections	~EOU*
MWD+IGRF	IGRF or WMM	-	+10%
MWD	BGGM or MVSD	-	Standard
MWD+HD	HDGM or MVHD	-	-10%
MWD+IFR1	Local In-Field Referencing (IFR1)	-	-30%
MWD+IFR1+MS	Local In-Field Referencing (IFR1)	Multi-Station	-50%

*Approximate values, actual EOU sizes depend on location and orientation of wellbore

New set of consolidated error models from the Operator Wellbore Survey Group. These are available as tool codes for well planning software.

Conclusions and Outlook

- Important to use annually updated models
- Don't mix models & degrees in secular variation corrections
- New line of global models
 - Standard Definition MVSD to degree 133 (tool code MWD)
 - High Definition MVHD to degree 1000 (tool code MWD+HD)
- IFR models have to cover the spectrum without gaps
 - Adding a crustal correction from a 80x80 km grid to a main field model can result in declination errors of over 1 degree
 - Ground shots: Still need a transform to give the declination at depth
- New aeromagnetic surveys in areas with strong crustal anomalies
 - Provide opportunity for testing and validation studies
 - Particularly interested in MWD versus Gyro comparisons

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